

## **Electronic Supplementary Materials**

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# **Effects of moisture content and dry bulk density on the thermal conductivity of compacted backfill soil**

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## Data S1 Sampling sites and characterization methods

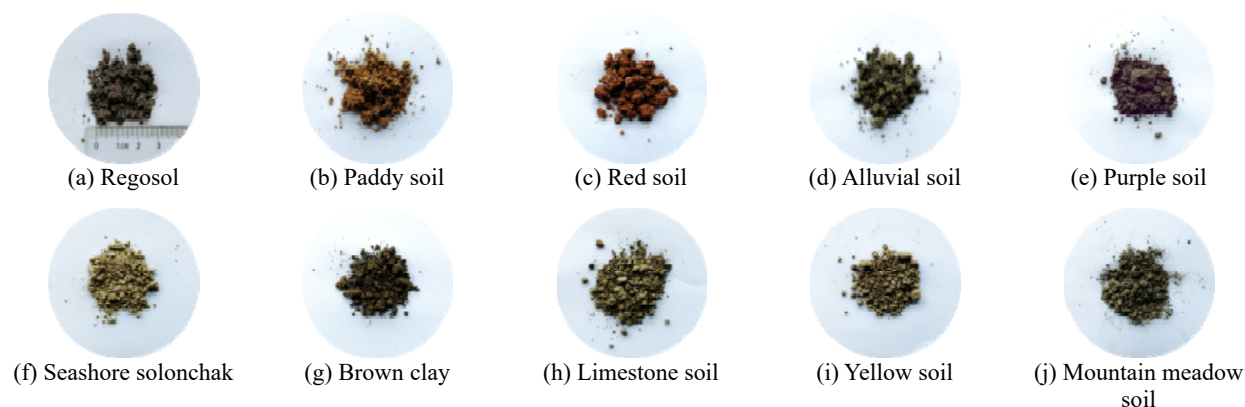
China has conducted two national soil censuses from 1958 to 1960 and from 1979 to 1985, which were generally based on townships and villages. On this basis, China's soil classification system has been gradually established. According to the soil classification standard in China (AQSIQ and SA, 2009), there are mainly ten typical soils in Zhejiang Province, including regosol, paddy soil, red soil, etcetera. Red soil accounts for about 70%, which is the most widely distributed (Wu et al., 2014).

Soils were sampled from ten different districts, counties or county-level cities in Zhejiang Province, including Anji, Lin'an, Yuhang, etcetera. The sampling sites were based on the second national soil

census (NSISP, 2019). Table S1 shows the specific differences of longitude and latitude between the second national soil census sites and the sampling sites. The current landform has undergone great changes compared to a few decades ago. In addition, soil sampling is required to avoid places that are difficult to excavate, such as paddy fields, mud pools and construction sites. Therefore, there are certain differences in longitude and latitude between the second national soil census sites and the sampling sites, and the absolute values of the differences do not exceed 0.04°. The appearance of the soil samples is shown in Fig. S1.

**Table S1 Comparison of longitude and latitude between the second national soil census sites and sampling sites**

Soil type	Region	The second national soil census site		Sampling site	
		East longitude (°)	North latitude (°)	East longitude (°)	North latitude (°)
Regosol	Anji	119.6240	30.5253	119.6117	30.5264
Paddy soil	Lin'an	119.7490	30.2528	119.7880	30.2712
Red soil	Yuhang	119.9100	30.2919	119.9080	30.2917
Alluvial soil	Shangyu	120.8681	30.0331	120.8683	30.0273
Purple soil	Fenghua	121.4069	29.6551	121.3860	29.6499
Seashore solonchak	Ninghai	121.4368	29.4791	121.4542	29.4507
Brown clay	Putuo	122.3955	29.9135	122.3837	29.9126
Limestone soil	Tonglu	119.6507	29.6798	119.6496	29.6857
Yellow soil	Wuyi	119.8163	28.8927	119.8063	28.9128
Mountain meadow soil	Liandu	119.7419	28.2958	119.7293	28.2927



**Fig. S1 Appearance of typical soil samples in Zhejiang Province, China**

To classify the soils by texture, the particle size distribution was characterized by a TM-85 soil densitometer, which is based on the principle of soil layered deposition in water according to particle size (MOHURD, 2019). To analyze the effect of

chemical composition on soil thermal conductivity, the mass percentage of minerals and organic matters were characterized respectively by X-ray diffraction method and dichromate titration-external heating method. X-ray diffraction was used to obtain the

diffraction pattern of soil samples by X-ray, followed by a comparison of the diffraction characteristics, such as peak type peak intensity and D-value, with the standard mineral types, so as to judge the mineral type and content in the samples (Xu et al., 2020). Dichromate titration-external heating method is to use excess potassium dichromate-sulfuric acid

solution to oxidize soil organic carbon under heating conditions, titrate the excess potassium dichromate with standard ferrous sulfate solution, and then calculate the organic matter content in soil samples according to the amount of potassium dichromate consumed and a series of correction coefficients (MOA, 2006).

## Data S2 Definition of compaction degree

Compaction degree is divided into absolute compaction degree  $C_{abs}$  and relative compaction degree  $C_{rel}$ , and its definition is based on dry bulk density  $\rho_d$ . Dry bulk density  $\rho_d$  refers to the ratio of the mass of the soil solid phase  $m_s$  to the total soil volume  $V$ , which reflects the compactness of the solid phase. The fully compacted dry bulk density  $\rho_{d,fc}$  of the soil is different at each moisture content. Soil samples could be fully compacted at each moisture content through the method of Proctor compaction test (MOHURD and AQSIQ, 2019). Then the relationship curve between fully compacted dry bulk density  $\rho_{d,fc}$  and moisture content  $\theta_m$  is drawn with  $\rho_{d,fc}$  as the ordinate and  $\theta_m$  as the abscissa, which is also called Proctor curve (Menaceur et al., 2021). The ordinate and abscissa of the peak point on the curve represent the maximum fully compacted dry bulk density  $\rho_{d,fc,max}$  and the optimal moisture content  $\theta_{m,opt}$  of the soil, respectively. Geotechnical engineering defines absolute compaction degree  $C_{abs}$  as the ratio of dry bulk density  $\rho_d$  to maximum fully compacted dry bulk density  $\rho_{d,fc,max}$  (MOHURD and AQSIQ, 2019). Similarly, relative compaction degree  $C_{rel}$  is defined as the ratio of dry bulk density  $\rho_d$  to fully compacted

dry bulk density at the current moisture content. The definition formulae of the vital parameters mentioned in this paragraph are shown in Eqs. (S1) ~ (S3).

$$\rho_d = \frac{m_s}{V} \quad (S1)$$

$$C_{abs} = \frac{\rho_d}{\rho_{d,fc,max}} \quad (S2)$$

$$C_{rel} = \frac{\rho_d}{\rho_{d,fc}} \quad (S3)$$

In order to make the research results significant to the engineering guiding, the relative compaction degree of soil samples was strictly controlled rather than the dry bulk density. Soil samples in this study were fully compacted by the provisions of the Proctor compaction test mentioned above. In other words, although the dry bulk density of each soil sample was different, the relative compaction degree reached 100%. Except for a few cases, the absolute compaction degree was greater than 85%, which basically met the provisions about the absolute compaction degree of subgrade backfill soil (MOT, 2019).

## Data S3 Apparatus and procedures of soil thermal conductivity measurement

The soil sample is filled in the cylindrical container, whose top cover has several position holes for probe insertion. The water bath case with a manageable temperature range 0~100 °C is used to control the temperature of the soil sample in the cylindrical container. The thermocouple connected with a temperature recorder is vertically inserted into the soil sample. The KD2 Pro thermal properties

analyzer is based on the thermal probe technique of the transient measurement method, whose major advantages is effectively minimizing the effect of moisture migration on the temperature distribution of the soil sample (Xu et al., 2019). The principles of the thermal probe technique are detailed in references (Modi et al., 2014; Kim et al., 2014).

The TR-1 probe with 100 mm in length and 2.4

mm in diameter is suit for measuring the thermal conductivity of porous materials. The range of thermal conductivity measured by the TR-1 probe is 0.1~4.0 W/(m·K), and the errors are  $\pm 0.02$  W/(m·K) at the range from 0.1 to 0.2 W/(m·K) and  $\pm 10\%$  that from 0.2 to 4.0 W/(m·K) (DDI, 2016). Note that the TR-1 probe should be fully contacted with soils since the thermal conductivity is calculated based on time and temperature of heat transfer process. In this sense, the length of the probe is a factor included in the thermal conductivity calculation of the KD2 Pro instrument. However, there is a stainless steel plate above the soil container as the top cover, which leads to a gap between the upper part of the probe and surface of the sample. The thickness of the plate is 1 mm, that is, the gap is only 1% of the length of the TR-1 probe. Therefore, compared to the systematic error ( $\pm 10\%$ ) mentioned above, the error caused by the stainless steel plate could be ignored. When the moisture content is relatively high, moisture exchange will occur between the sample and the surroundings, which may cause an error even greater than the systematic error. In view of this, the stainless steel plate was utilized as the top cover to seal the cylindrical container and avoid moisture exchange.

The preparation of soil samples was carried out according to the steps as follows. (1) The soil was dried in an incubator at 105 °C for 8~12 h to evaporate all the natural water it contained. (2) The soil was ground thoroughly with a mortar. (3) Based on the target moisture content, a certain quality of water was added to the soil and then mixed evenly. (4) According to the provisions of Proctor compaction test, the soil was fully compacted with a compactor composed of a compaction casing with an

inner diameter of 102 mm, a pile casing matching with the compaction casing and a cast iron hammer with a mass of 2.5 kg. Specifically, the soil in the compaction casing was compacted in three layers, and each layer was hammered 25 times. The hammer fell down freely from a height of 305 mm for each blow.

The measurement of soil thermal conductivity was carried out according to the steps as follows. (1) The soil sample was moved from the compaction casing into the cylindrical container carefully to avoid damaging its structure and changing its compaction degree. (2) The top cover was connected with the cylindrical container through the flange structure. (3) The TR-1 probe matched with the KD2 Pro and the thermocouple connected with the temperature recorder were both inserted into the soil sample through the position holes. (4) Waterproof tape was used to cover the gaps of apparatus connection to prevent moisture exchange, and also fix the position of the TR-1 probe and the thermocouple. (5) The container was placed in the water bath case and the temperature was set to the target value of 20 °C. When the soil temperature measured by the thermocouple had been stable near the target temperature for more than 1 h, the thermal conductivity measurement could be carried out. (6) The time of each measurement was 5 minutes, and the interval between two adjacent measurements was supposed to be more than 15 minutes to ensure sufficient cooling of the probe, so as to reduce the measurement errors. In addition, errors were also reduced by measuring repeatedly, eliminating bad values, and changing the position hole where the TR-1 probe inserted in for each measurement.

## Data S4 Original measurement data

**Table S2 Original measurement data of density and thermal conductivity of soil samples**

Soil type	$\theta_m$ (%)	$\rho_w$ (g/cm <sup>3</sup> )	$\rho_d$ (g/cm <sup>3</sup> )	$k$ (W/(m·K))						
				Measurement (five times)					Mean	Standard deviation
Regosol	0	1.366	1.366	0.218	0.213	0.209	0.208	0.209	0.211	0.004
	5	1.509	1.437	0.582	0.630	0.637	0.669	0.624	0.628	0.031
	10	1.661	1.510	1.028	1.108	1.063	0.987	1.180	1.073	0.074
	15	1.861	1.618	1.564	1.702	1.693	1.839	1.740	1.708	0.099
	20	2.043	1.703	1.943	2.189	2.108	1.879	2.135	2.051	0.133
	25	2.049	1.639	2.002	1.790	1.947	1.981	1.856	1.915	0.090
Paddy soil	0	1.326	1.326	0.208	0.216	0.209	0.213	0.214	0.212	0.003
	5	1.429	1.361	0.514	0.513	0.487	0.508	0.525	0.509	0.014
	10	1.547	1.406	0.943	0.850	0.941	0.908	0.858	0.900	0.044
	15	1.681	1.462	1.240	1.403	1.378	1.389	1.420	1.366	0.072
	20	1.810	1.508	1.631	1.787	1.804	1.634	1.710	1.713	0.082
	25	1.890	1.512	1.847	2.082	1.962	1.804	1.856	1.910	0.112
Red soil	0	1.314	1.314	0.202	0.210	0.209	0.204	0.200	0.205	0.004
	5	1.460	1.391	0.533	0.548	0.527	0.540	0.563	0.542	0.014
	10	1.540	1.400	0.898	0.885	0.820	0.810	0.907	0.864	0.046
	15	1.629	1.417	1.112	1.058	1.007	1.061	1.124	1.072	0.047
	20	1.717	1.431	1.301	1.241	1.425	1.396	1.355	1.344	0.074
	25	1.855	1.484	1.904	1.889	1.878	1.796	1.825	1.858	0.046
Alluvial soil	0	1.379	1.379	0.242	0.247	0.250	0.252	0.248	0.248	0.004
	5	1.497	1.425	0.689	0.659	0.677	0.614	0.650	0.658	0.029
	10	1.616	1.469	1.108	1.057	1.026	0.970	1.020	1.036	0.051
	15	1.800	1.565	1.413	1.435	1.500	1.340	1.379	1.413	0.060
	20	1.945	1.621	1.992	1.780	1.980	1.839	1.843	1.887	0.094
	25	2.004	1.603	1.801	1.896	1.780	1.650	1.826	1.791	0.090
Purple soil	0	1.392	1.392	0.244	0.246	0.250	0.244	0.241	0.245	0.003
	5	1.492	1.421	0.596	0.625	0.674	0.598	0.632	0.625	0.032
	10	1.569	1.427	0.985	0.938	1.031	1.017	0.940	0.982	0.043
	15	1.655	1.439	1.217	1.224	1.116	1.307	1.109	1.195	0.083
	20	1.792	1.493	1.391	1.735	1.520	1.458	1.552	1.531	0.130
	25	1.875	1.500	1.927	1.820	1.832	1.901	1.764	1.849	0.065
Seashore solonchak	0	1.494	1.494	0.242	0.243	0.246	0.248	0.246	0.245	0.002
	5	1.603	1.527	0.680	0.756	0.790	0.743	0.787	0.751	0.045
	10	1.715	1.559	1.017	1.168	1.107	1.134	1.210	1.127	0.073
	15	1.828	1.590	1.493	1.405	1.610	1.385	1.562	1.491	0.097
	20	2.001	1.667	1.836	1.961	1.998	1.703	1.776	1.855	0.124
	25	2.047	1.638	1.905	1.885	1.784	1.816	1.922	1.862	0.060
Brown clay	0	1.325	1.325	0.206	0.210	0.215	0.208	0.215	0.211	0.004
	5	1.481	1.411	0.640	0.735	0.702	0.648	0.682	0.681	0.039
	10	1.583	1.439	0.928	0.940	0.857	0.923	0.926	0.915	0.033
	15	1.700	1.479	1.268	1.472	1.469	1.402	1.373	1.397	0.084
	20	1.888	1.574	1.524	1.644	1.671	1.596	1.623	1.612	0.056
	25	1.943	1.554	1.644	1.588	1.584	1.642	1.538	1.599	0.045
Limestone soil	0	1.411	1.411	0.251	0.247	0.253	0.252	0.254	0.251	0.003
	5	1.504	1.433	0.701	0.713	0.770	0.662	0.742	0.718	0.041
	10	1.592	1.448	1.043	1.205	1.187	1.247	1.135	1.163	0.078
	15	1.688	1.468	1.792	1.587	1.690	1.646	1.800	1.703	0.092
	20	1.910	1.592	2.028	2.019	1.912	1.947	2.004	1.982	0.050

	25	1.912	1.529	1.966	1.907	1.987	1.942	1.985	1.957	0.034
Yellow soil	0	1.270	1.270	0.213	0.206	0.207	0.212	0.210	0.210	0.003
	5	1.340	1.277	0.273	0.294	0.285	0.279	0.284	0.283	0.008
	10	1.392	1.265	0.424	0.392	0.385	0.417	0.412	0.406	0.017
	15	1.505	1.309	0.737	0.808	0.658	0.707	0.765	0.735	0.057
	20	1.584	1.320	0.872	0.892	0.830	0.943	0.915	0.890	0.043
	25	1.620	1.296	1.141	0.958	1.044	1.063	1.018	1.045	0.067
Mountain meadow soil	0	1.462	1.462	0.280	0.276	0.279	0.286	0.282	0.281	0.004
	5	1.567	1.492	0.706	0.732	0.603	0.666	0.631	0.668	0.053
	10	1.657	1.507	1.140	0.977	1.131	1.073	0.985	1.061	0.078
	15	1.742	1.515	1.463	1.636	1.692	1.698	1.596	1.617	0.096
	20	1.844	1.537	1.707	2.012	1.906	1.784	1.885	1.859	0.117
	25	1.919	1.535	1.975	2.036	1.913	1.987	1.965	1.975	0.044

## Data S5 Procedures and results of regression analysis

First, a correlation analysis was adopted to screen out the main parameters affecting soil thermal conductivity. Under the condition that the significance level  $\alpha$  equals to 0.05, the correlation analysis results between various parameters and the thermal conductivity of fully compacted soils are shown in Table S3. The coefficient R is the Pearson product moment correlation coefficient. It is obvious that the moisture content and dry bulk density have a strong positive correlation with soil thermal conductivity, while texture and chemical composition have a relatively weak correlation with soil thermal conductivity. In addition, texture and chemical composition are the natural properties of soils and are not easy to change, while moisture content and dry bulk density are easy to change by external conditions.

Then, on the basis of the above correlation analysis, a regression analysis based on moisture content  $\theta_m$  and dry bulk density  $\rho_d$  as the independent parameters was adopted to summarize the prediction formulae. It should be noted that, due to the discrepancies of thermal conductivity among various types of soils, it is more meaningful to propose the prediction formula for the thermal conductivity of a certain category rather than for each specific soil for engineering involving soil backfilling. Thus, formulae for two categories of fully compacted soils were summarized. The formula for clay loam and loam was based on all the data of paddy soil, red soil, purple soil, seashore solonchak, brown clay and mountain meadow soil, while that for sandy loam and loamy sand was based on all the data of regosol, alluvial soil and limestone soil. The data of yellow soil was not involved in the regression analysis because of the huge discrepancy between its chemical composition and thermal conductivity with the other nine typical soils.

It is considered that the confidence of the model reaches 95% when the P-value of all the regression coefficients in the fitting formula are less than 0.05. However, in the fitting formulae obtained by binary linear regression with  $\theta_m$  and  $\rho_d$ , the P-value of some regression coefficients is slightly greater than 0.05. Thus, the final formulae in Table S3 are based on the binary linear regression with  $\theta_m^{1/2}$  and  $\rho_d^{1/2}$ , that meet

the confidence requirements and have a coefficient of complex determination  $R^2$  exceeds 0.9. The formulae in Table S3 are applicable to compacted backfill soils which have the similar texture classification and chemical composition to the typical soils in this study. Other results of the binary linear regression analysis are shown in Table S4.

**Table S3 Results of correlation analysis between various parameters and soil thermal conductivity**

Parameters affecting soil thermal conductivity	R
Moisture content	0.914
Dry bulk density	0.786
Mass percentage of clay	-0.079
Mass percentage of silt	-0.105
Mass percentage of sand	0.110
Mass percentage of quartz	0.225
Mass percentage of other minerals	-0.253
Mass percentage of organic matters	0.139

**Table S4 Results of binary linear regression analysis**

(i) Clay loam &amp; loam (6 types)

(a)

	Freedom	Regression sum of square	Mean square deviation	F	Significance F
Regression analysis	2	11.180	5.590	217.704	0.000
Residual error	33	0.847	0.026	---	---
Sum up	35	12.027	---	---	---

(b)

	Coefficients	Standard deviation	t-Stat	P-value	Confidence interval	
					Lower 95%	Upper 95%
Intercept	-4.342	1.211	-3.584	0.001	-6.806	-1.877
Variable 1 ( $\theta_m^{1/2}$ )	2.821	0.210	13.412	0.000	2.393	3.249
Variable 2 ( $\rho_d^{1/2}$ )	3.773	1.032	3.654	0.001	1.672	5.873

(ii) Sandy loam &amp; loamy sand (3 types)

(a)

	Freedom	Regression sum of square	Mean square deviation	F	Significance F
Regression analysis	2	7.015	3.507	122.702	0.000
Residual error	15	0.429	0.029	---	---
Sum up	17	7.444	---	---	---

(b)

	Coefficients	Standard deviation	t-Stat	P-value	Confidence interval	
					Lower 95%	Upper 95%
Intercept	-5.698	2.222	-2.564	0.022	-10.435	-0.962
Variable 1 ( $\theta_m^{1/2}$ )	2.718	0.453	5.994	0.000	1.751	3.684
Variable 2 ( $\rho_d^{1/2}$ )	4.958	1.905	2.603	0.020	0.899	9.017

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